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EVOLUTION OF THE EARTH'S ATMOSPHERE

"We living things are a late outgrowth of the metabolism of our Galaxy. The carbon that enters so importantly into our composition was cooked in the remote past in a dying star. From it at lower temperatures nitrogen and oxygen were formed. These, our indispensable elements, were spewed out into space in the exhalations of red giants and such stellar catastrophes as supernovae, there to be mixed with hydrogen, to form eventually the substance of the sun and planets, and ourselves. The waters of ancient seas set the pattern of ions in our blood. The ancient atmospheres molded our metabolism."

---George Wald

nitrogen, 21% oxygen and 1% argon, with traces of carbon dioxide, water vapor and ozone. The atmospheres of Mars and Venus, on the other hand, are predominantly composed of carbon dioxide, while those of Jupiter and Saturn contain mainly hydrogen, helium, methane and ammonia. Such a wide variety in the composition of the atmospheres of the planets appears most intriguing when one considers that all nine planets were probably formed at the same time and out of the same chemically homogeneous mixture of gas and dust; that is, the primitive solar nebula:

The most likely explanation for this diversity in composition seems to be that the planetary atmospheres have undergone

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important evolutionary changes during their long history of about 4.5 billion years. On the earth, the evidence for such changes is manifold. Ever since Louis Pasteur demonstrated that spontaneous generation of life does not take place at the present time, it has become obvious that if life originated on the earth, as it probably did, the atmospheric conditions in the early history of the planet must have been quite different in order for the chemical evolution to have taken place. Also, it appears that relative to the sun, the earth is deficient by several orders of magnitude, not only in hydrogen and helium, which being the lighest can be assumed to have escaped from the gravitational field of the earth, but also in carbon, nitrogen, and the cosmically abundant rare gases; namely, Ne, A³⁶, Kr and Xe, (Table 1). comparison of the deficiency factors of these elements with those of non-volatiles like Na, Mg, Al and Si, strongly suggests that the earth lost all elements which could have been gaseous at temperatures of a few hundred degrees during the very early history of the planet. The relatively small amounts of carbon, hydrogen, nitrogen and oxygen now present in the crust, hydrosphere and atmosphere were probably outgassed from the interior and accumulated slowly during the lifetime of the earth.

Both biochemical arguments and geological evidence, therefore, strongly suggest that the atmosphere of the earth has undergone major evolutionary changes during its long history. Whether the atmospheres of Mars and Venus have also gone through similar stages of evolution is a question which cannot be answered without more knowledge about their composition. Future in situ exploration of these planets and a search for the rare gases, especially Ne, which is deficient on earth by a factor of 10^{10} , will provide important clues to the origin of these atmospheres.

For the earth, there is convincing evidence that the present atmosphere and hydrosphere arose largely from the interior via volcanic emanations. But the sequence of events which led to present-day composition of N₂ and O₂ has yet to be established. What is the history of the volatiles now present at the surface of the earth? Has the C, N, O and H always been in the form of CO₂, N₂, and H₂O and H₂, or did carbon and nitrogen combine with hydrogen early in the earth's history to form CH₄ and NH₃? Under what atmospheric conditions did life originate on earth and how did the appearance of life change the atmosphere? These are some of the basic questions which must be answered in order to paint a coherent picture of the evolution of the earth's atmosphere.

Opinions on these questions are many and varied; sometimes they are almost diametrically opposed. The Oparin-Urey theory of the origin of life on the earth, supported by the laboratory experiments of Miller and more recently of Ponnamperuma, suggest a primitive atmosphere composed mainly of CH₄ with small amounts of NH₃, H₂ and H₂O vapor. On the other hand, the school of thought led by Abelson, and supported also by laboratory experiments on the synthesis of amino acids, holds that the early atmosphere of the earth was made up of CO₂, CO, H₂, N₂ and H₂O vapor.

Geologists are also divided on the subject. Holland has presented a model for the evolution of the atmosphere in which, during the very early stage after the formation of the earth and at the commencement of outgassing, the dominant gases could have been CH_4 and H_2 , provided free hydrogen did not 'escape' as rapidly as it does today. Rubey, on the other hand, believes that the early atmosphere was probably made of CO_2 and N_2 because not enough hydrogen was available to keep CH_4 from converting into CO_2 . Holland's model is supported by the calculations of Rasool and McGovern, who have investigated the thermal properties of model primitive atmospheres of the earth. They a find that $\operatorname{in/99\%} \operatorname{CH}_4 - \operatorname{1\%} \operatorname{H}_2$ atmosphere the average exospheric

temperature may be as low as 650° K (cf. present day value 1500° K), making the escape of hydrogen a relatively slow phenomenon. However, Abelson has argued that if methane were abundant in the primitive atmosphere, the earliest rocks should contain unusual amounts of organic matter, which apparently is not the case.

Despite the disagreement over the composition of the primitive atmosphere, it is almost certain that it was devoid of free oxygen. How and when did free oxygen then become a major constituent of the atmosphere? The late Lloyd V. Berkner and L.C. Marshall have presented detailed calculations of the photochemistry of H₂O, indicating that free oxygen was limited to \sim 0.1% of the present atmospheric level during the entire prebiological history, accumulating slowly to the present amount since the start of photosynthesis about 800 million years ago. This is probably the most widely accepted explanation for the growth of oxygen in the earth's atmosphere. However, difficulties arise when one attempts to construct an evolutionary model of the atmosphere which would be consistent from the prebiological period to the present. For example, it is not well understood whether the small amount of O, suggested to be present in the primitive atmosphere would not be sufficient to rapidly oxidize CH_4 and

 ${
m NH}_3$ into ${
m CO}_2$ and ${
m N}_2$. It may be that the production of even this small amount of free oxygen in the upper atmosphere by the photodissociation of water vaopr was inhibited by the presence of ${
m NH}_3$ itself.

Some of these important questions will be discussed in the forthcoming interdisciplinary symposium on the Evolution of the Earth's Atmosphere to be held during the Annual Meetings of the AAAS in New York City, under the chairmanship of Harrison Brown of the California Institute of Technology.

Because of the diversity of disciplines involved, an attempt has been made to bring together geologists, biologists, meteorologists and astrophysicists, with the hope of obtaining fresh ideas and a broader perspective on the subject. The program includes the following presentations: "Nature of the Early Atmosphere and Ocean" (P. Cloud), "Gravitational Escape of Hydrogen from Model Primitive Atmospheres" (S.I. Rasool and W.E. McGovern), "Simulated Atmospheres and Models of the Synthesis of the First Organisms" (S.W. Fox), "Evolution of N₂ - O₂ Atmosphere on Earth" (L.C. Marshall), Atmospheric CO₂, Past and Present" (H.D. Holland), followed by "Summary and Critique" (H. Brown). There is a chance that this discussion may bring

us a step closer to understanding the nature of the ancient atmospheres that molded our metabolism.

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TABLE 1.

(ATOMS/10,000 ATOMS Si)

0	WHOLE EARTH*	SOLAR SYSTEM**	DEFICIENCY FACTOR
	(a)	(b)	log (b/a)
Н	250	2.6 x 10 ⁸	6.0
Не	3.5×10^{-7}	2.1×10^{7}	13.8
С	14	135,000	4.0
N	0.21	24,400	5.1
0	35,000	236,000	0.8
Ne	1.2 x 10 ⁻⁶	23,000	10.3
Na	460	632	~ 0
Mg	8,900	10,500	~ 0
Al	940	851	~ 0
Si	10,000	10,000	0
A ³⁶	5.9×10^{-4}	2,280	6/ . 6
Kr	6 × 10 ⁻⁸	0.69	7.1
Xe	5 x 10 ⁻⁹	0.07	7.1

^{*} After Rubey, Ringwood and Mason

^{**} After Cameron